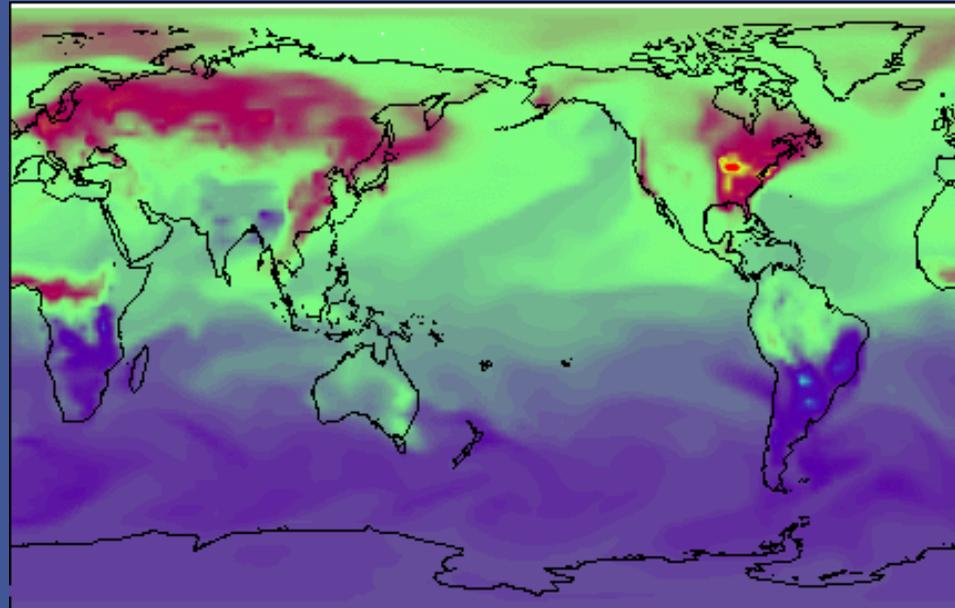


Offline transport models and the carbon cycle



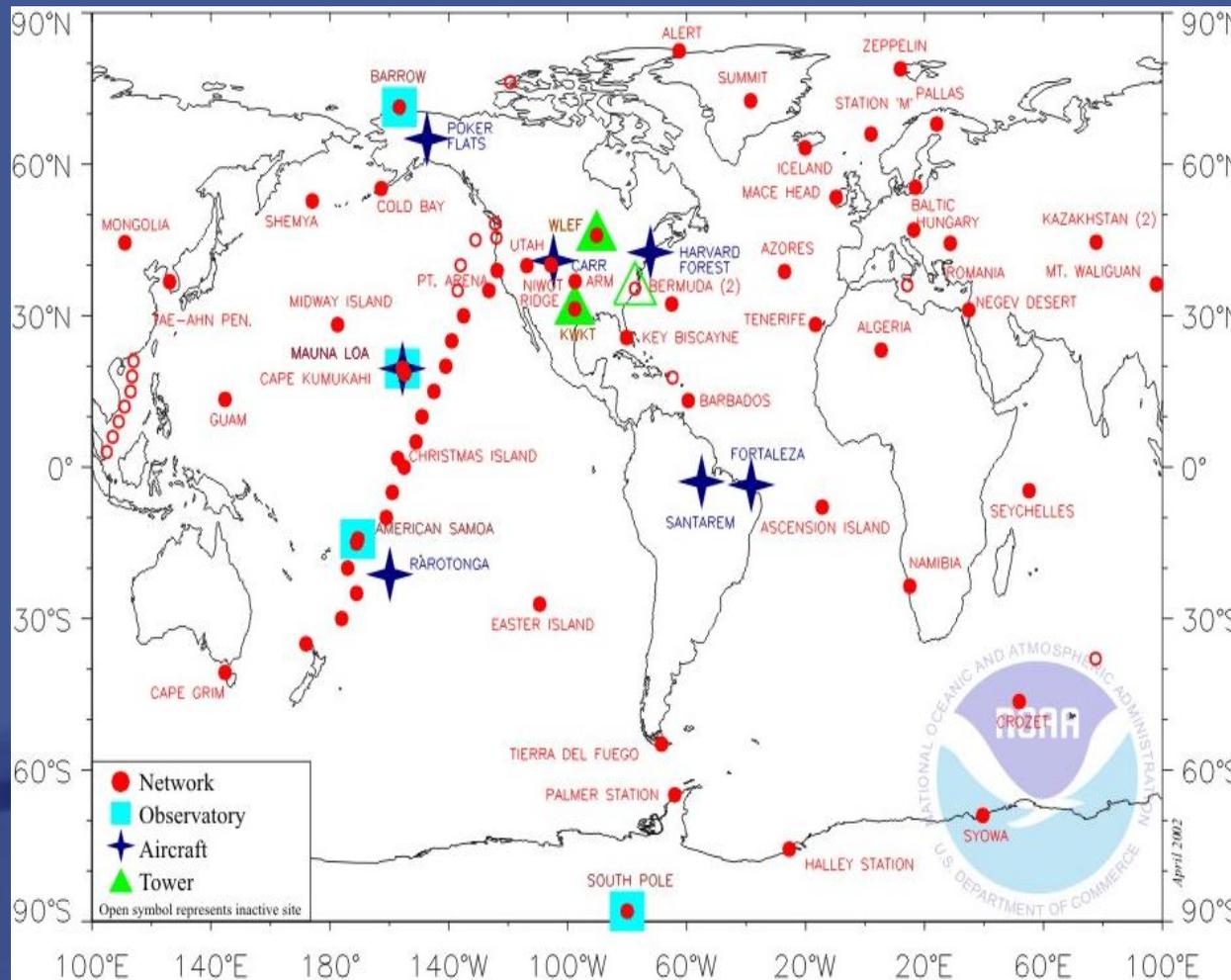
CO₂ Concentrations at the Surface



Scientific Questions

- Examine how atmospheric carbon concentrations have changed due to anthropogenic activities.
- Understand what and how the carbon budget affects the climate and thus the environment.

Schematic CO₂ Cycle



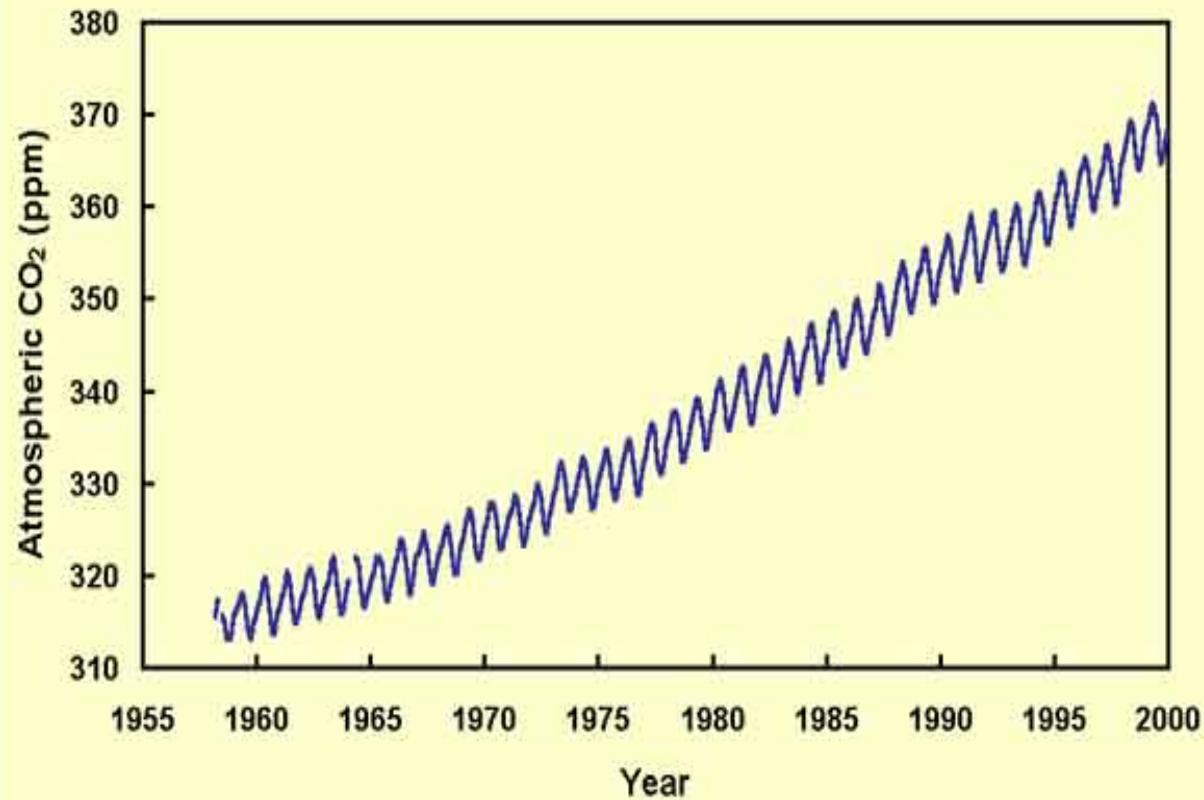
Observation network

Concentrations

Fluxes

Schematic CO₂ Cycle

Atmospheric CO₂ Concentration, Mauna Loa, HI (1958-1999)

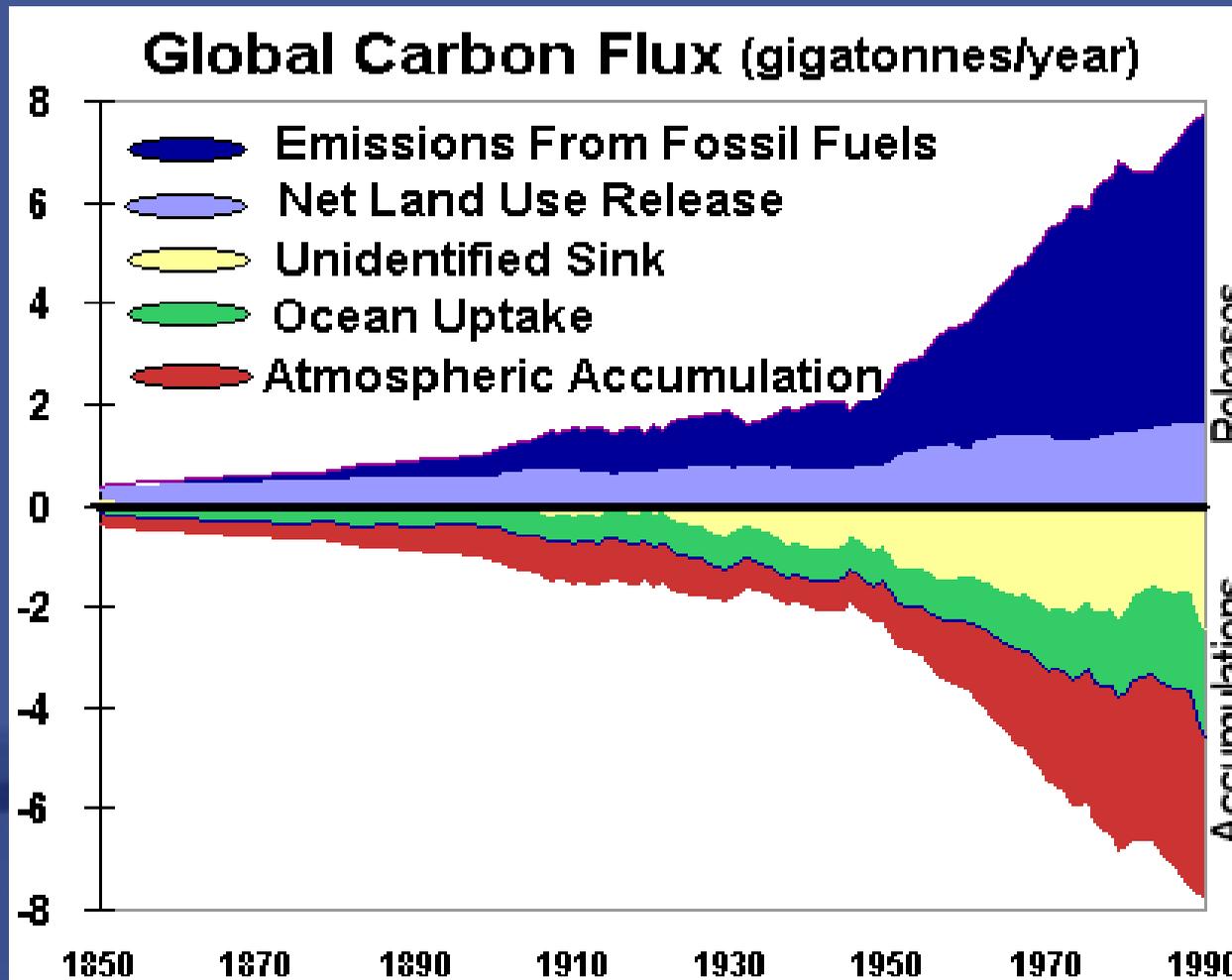


Observation network

Concentrations

Fluxes

Schematic CO₂ Cycle



Observation network

Concentrations

Fluxes

Transport Model

A (parametrized chemistry) transport model numerically solves the **constituent continuity equation** over the entire globe.

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- Carbons
- Ozone
- Aerosols
- Liquid water

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Continuity equation

$$\frac{\partial x}{\partial t} = -\mathbf{v}\nabla x + P(x) - L(x)$$

x : constituent

\mathbf{v} : velocity

$P(x), L(x)$: production and loss

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x : constituent

\mathbf{v} : velocity

$P(x), L(x)$: production and loss

⇒ where do we get \mathbf{v} from?

⇒ how to we solve the equation?

General Circulation Model

General circulation models are a **numerical representation** of the atmosphere and its phenomena over the entire Earth.

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Numerical representation

- Physical laws of flow
- Simplification
- Discretization

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Numerical representation

- Physical laws of flow
 - Simplification
 - Discretization
- Conservation of momentum
 - Conservation of mass
 - Equation of state for ideal gases
 - Conservation of energy
 - Conservation equation for water mass

General Circulation Model

General circulation models are a **numerical representation** of the atmosphere and its phenomena over the entire Earth.

Numerical representation

- Physical laws of flow
- Simplification
- Discretization

$$\frac{d\mathbf{v}}{dt} = -\alpha \nabla p - \nabla \phi + \mathbf{F} - 2\boldsymbol{\Omega} \times \mathbf{v}$$

$$\frac{\partial \rho}{\partial t} = -\nabla \cdot (\rho \mathbf{v})$$

$$p\alpha = RT$$

$$Q = C_p \frac{dT}{dt} - \alpha \frac{dp}{dt}$$

$$\frac{\partial \rho q}{\partial t} = -\nabla \cdot (\rho \mathbf{v} q) + \rho(E - C)$$

General Circulation Model

General circulation models are a **numerical representation** of the atmosphere and its phenomena over the entire Earth.

Numerical representation

- Physical laws of flow
 - **Simplification**
 - Discretization
- Spherical coordinates
 - Traditional approximation
 - Hydrostatic model
 - Quasi-geostrophic model
 - Barotropic model

General Circulation Model

General circulation models are a **numerical representation** of the atmosphere and its phenomena over the entire Earth.

Numerical representation

- Physical laws of flow
- Simplification
- **Discretization**
 - Leapfrog
 - Implicit (semi/fully)
 - Euler (forward/backward)

General Circulation Model

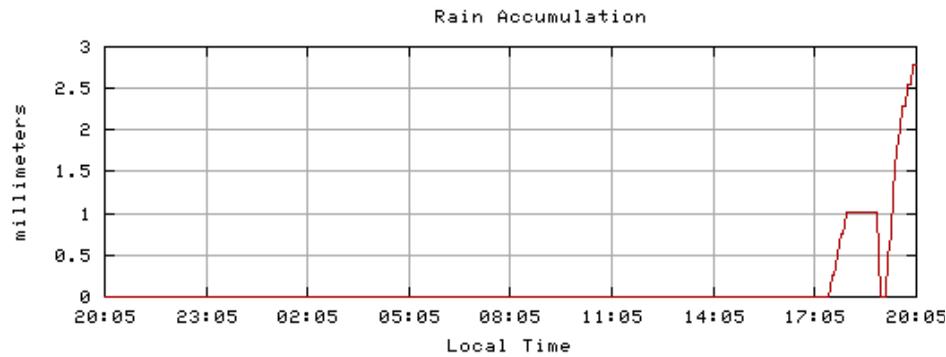
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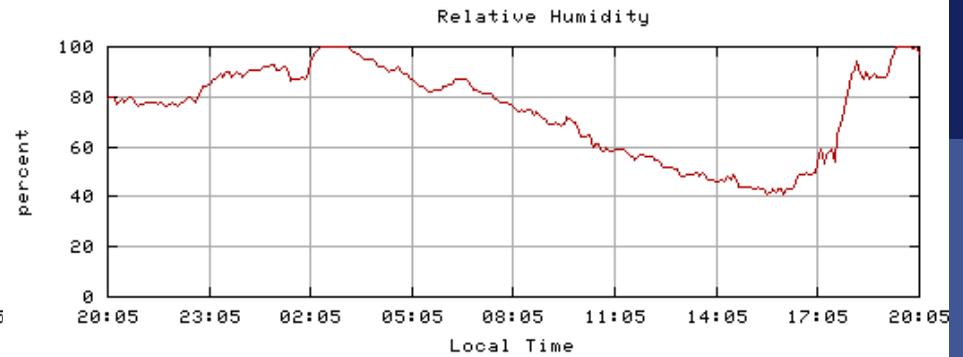
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Observations: Initial/boundary cond.

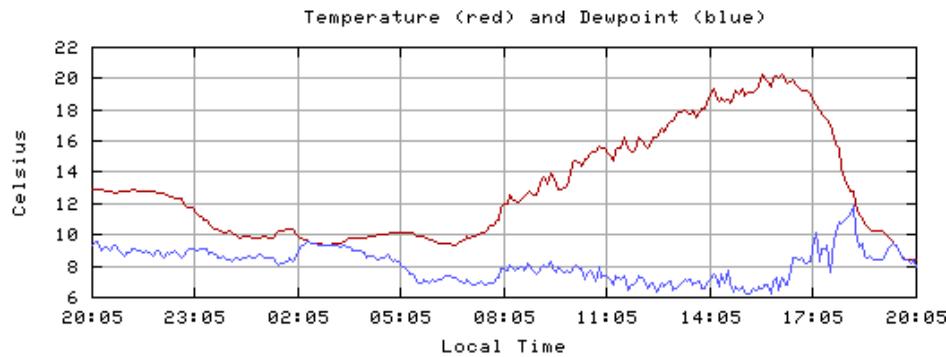
- Surface (towers, ships)
- Altitude (planes, balloons)
- Radar
- Satellites



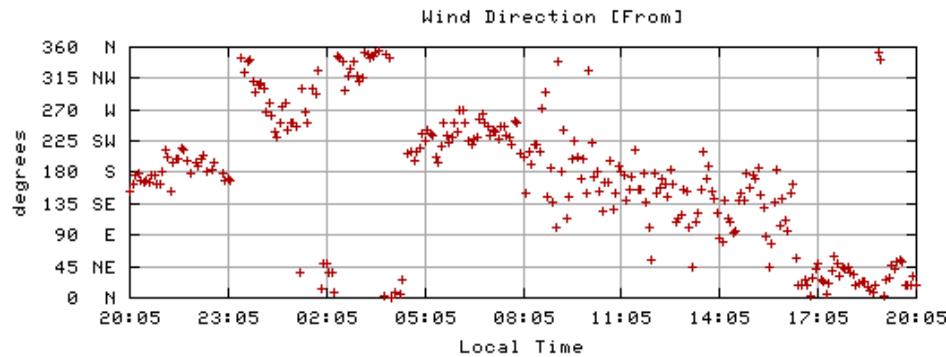
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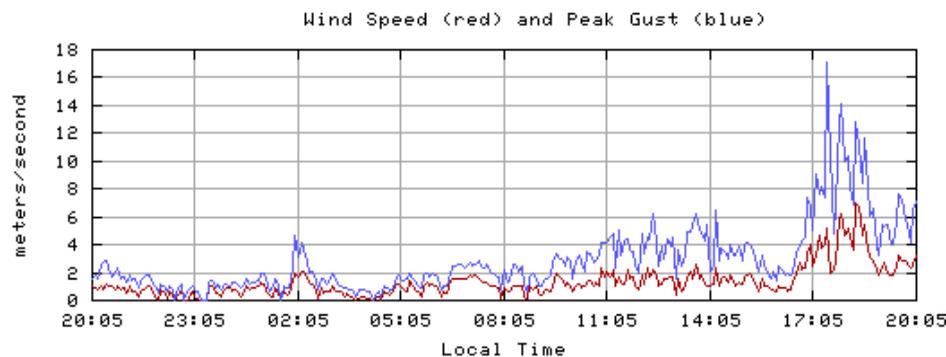
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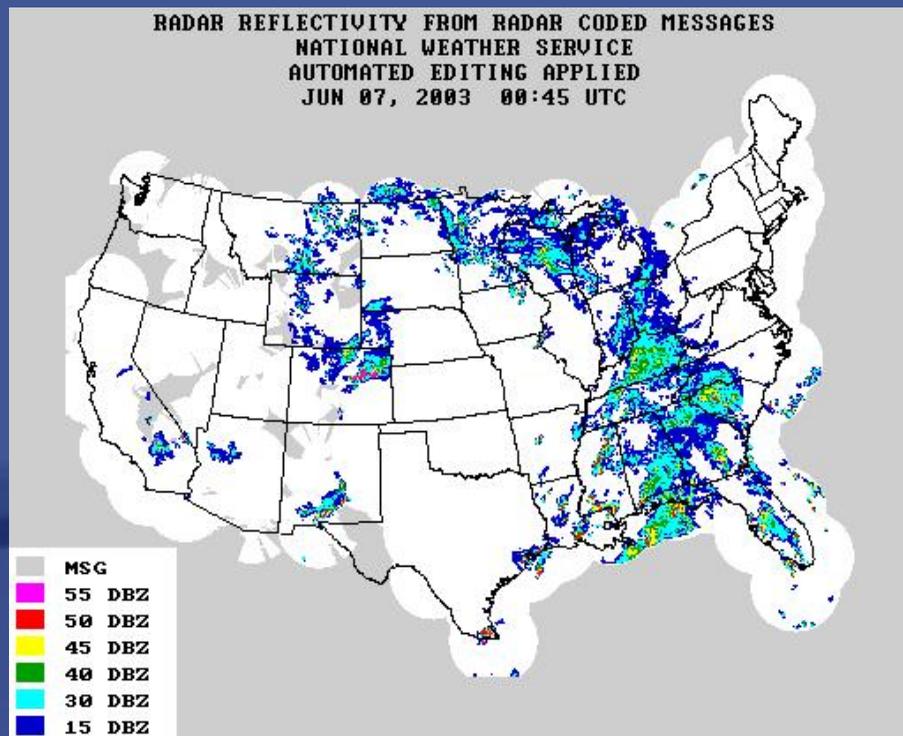
Monitors over the entire Earth.

Observations: Initial/boundary cond.

- Surface (towers, ships)
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General Circulation Model

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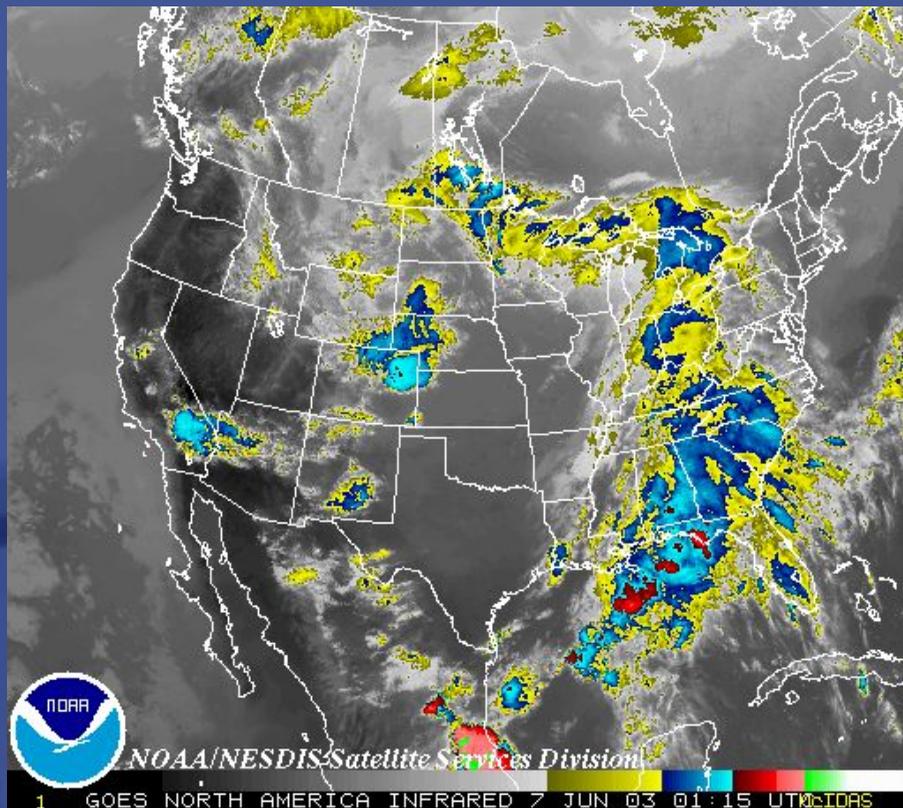


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Model output

Wind, humidity, temperature, pressure and specific density (at a “regular” grid).

Observations: Initial/boundary cond.

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- Altitude (planes, balloons)
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Offline Transport Model

In an offline transport model, the dynamics of the model are supplied by archived meteorological datasets.

Dynamics = winds, temperatures, surface pressures and cloud mass fluxes.

Modeling the CO₂ Cycle

We model the CO₂ cycle with:

- \mathbf{z}_t measured concentrations at a given network,
- \mathbf{x}_t actual concentrations (mixing ratios),
are dynamically constrained to $\mathbf{x}_{t-\delta}$ and fluxes,
- \mathbf{u}_t surface CO₂ fluxes,
represent the sources and sinks in the model.

We solve for the unknown fluxes $\{\mathbf{u}_t\}$

Inverse Method 1 2

The CO₂ surface flux problem is formulated as an optimality problem (simplified vector notation)

$$\begin{aligned} \min_{\{\mathbf{u}\}} & \sum_t (h(\mathbf{x}_t) - \mathbf{z}_t)^T \mathbf{W}_1 (h(\mathbf{x}_t) - \mathbf{z}_t) \\ & + \sum_t (\mathbf{u}_t - \mathbf{u}_t^{\text{con}})^T \mathbf{W}_2 (\mathbf{u}_t - \mathbf{u}_t^{\text{con}}) \\ & + (\mathbf{x}_0 - \mathbf{x}_0^{\text{con}})^T \mathbf{W}_3 (\mathbf{x}_0 - \mathbf{x}_0^{\text{con}}) \end{aligned}$$

subject to dynamical constraints

where:

$h(\cdot)$ is the measurement function,
 \mathbf{W}_i are weight matrices,
superscript 'con' are constraints,
subscript '0' are initial conditions.

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Inverse Method 1 2

Run model forward:

with initial concentrations and forced by a priori fluxes
↪ obtain modeled measurement history.

Run (adjoint) model backward:

forced by the weighted measurement differences
↪ get the adjoint to the concentrations,
↪ get new estimates for the fluxes and
initial concentrations.

Repeat until convergence is achieved.

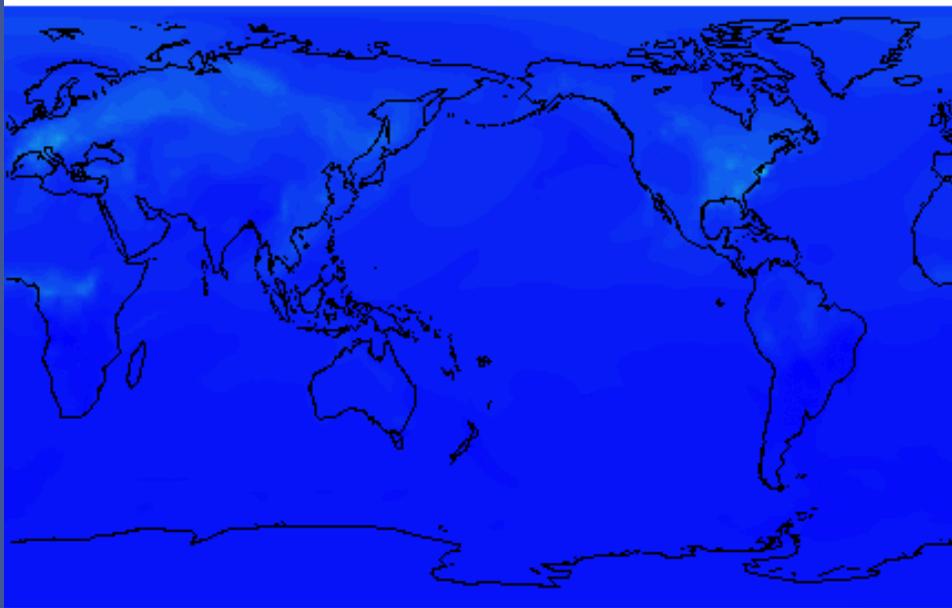
Statistical Work

Quantify the uncertainty of carbon sources and sinks on a regional scale.

- Acquaint with transport model
- First quantitative result of a sensitivity analysis
- Setup Gibbs sampler
-



Sensitivity Analysis: Simulation Setup

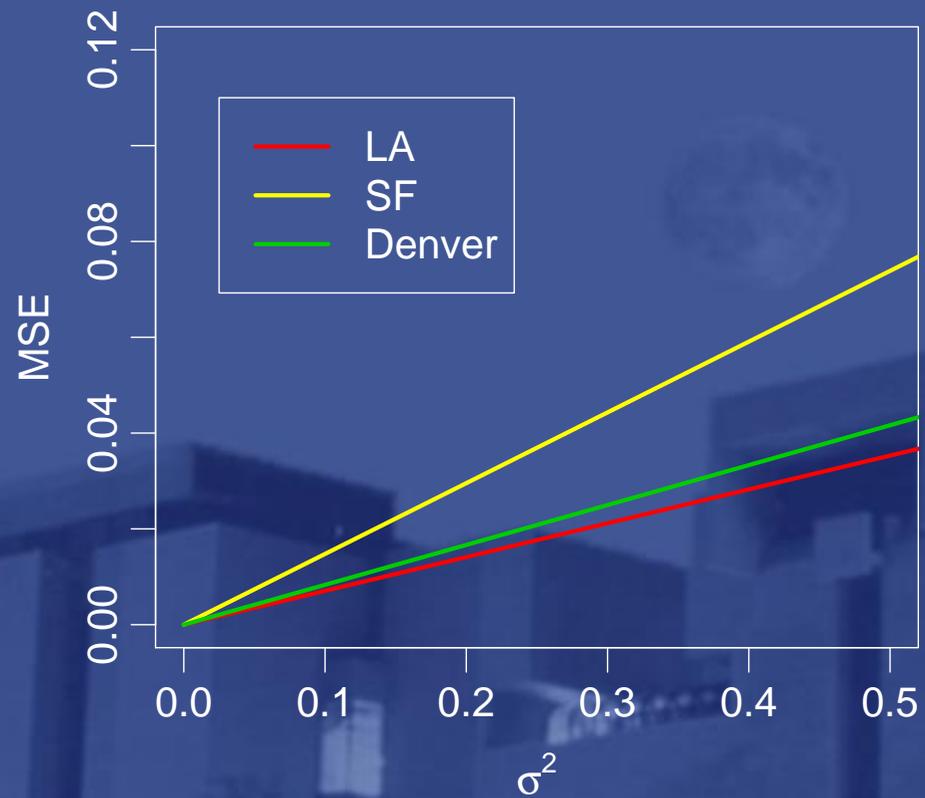


Transport model from NASA,
Goddard Space Flight Center.
PCTM, Lin/Rood advection code
w/ simple vertical mixing schemes.

- Eliminate sources and sinks
- Inject CO₂ “Dirac” point sources at LA, SF and Denver
- Run model forward with measurement and/or transport errors
- Try to recover the intensity with regression techniques after 4 days

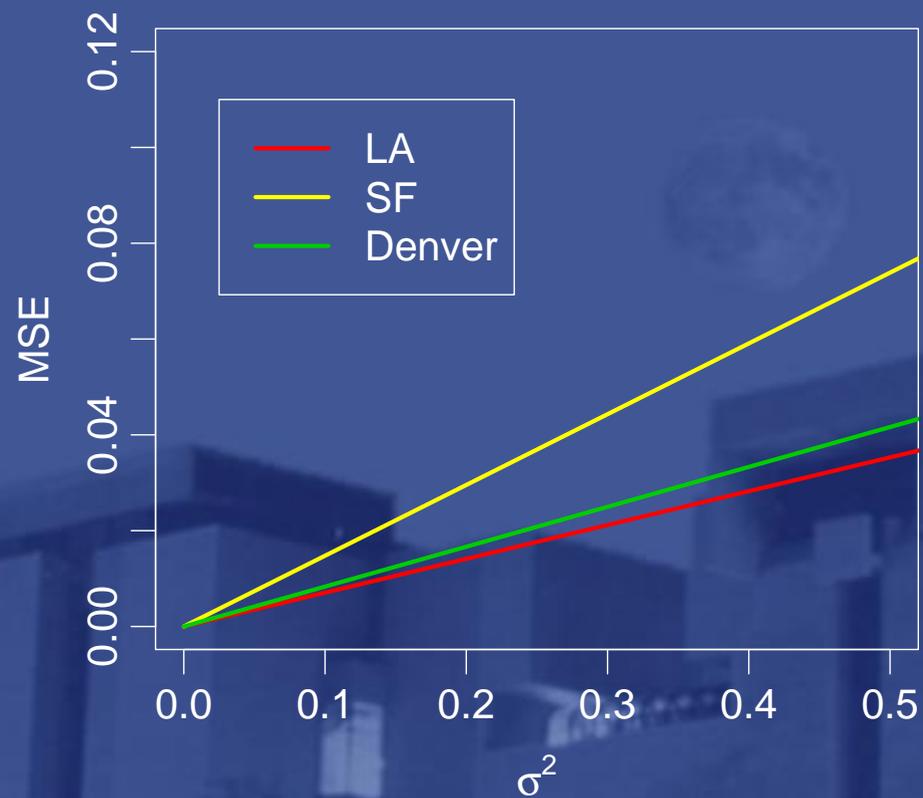
Sensitivity Analysis: Results

Measurement error

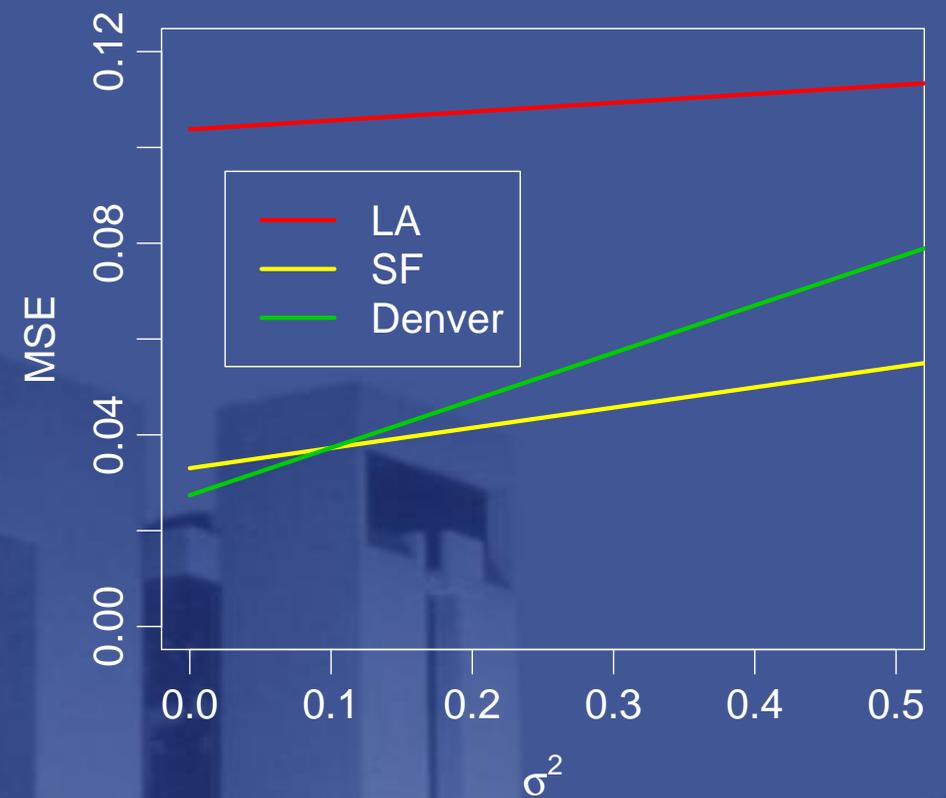


Sensitivity Analysis: Results

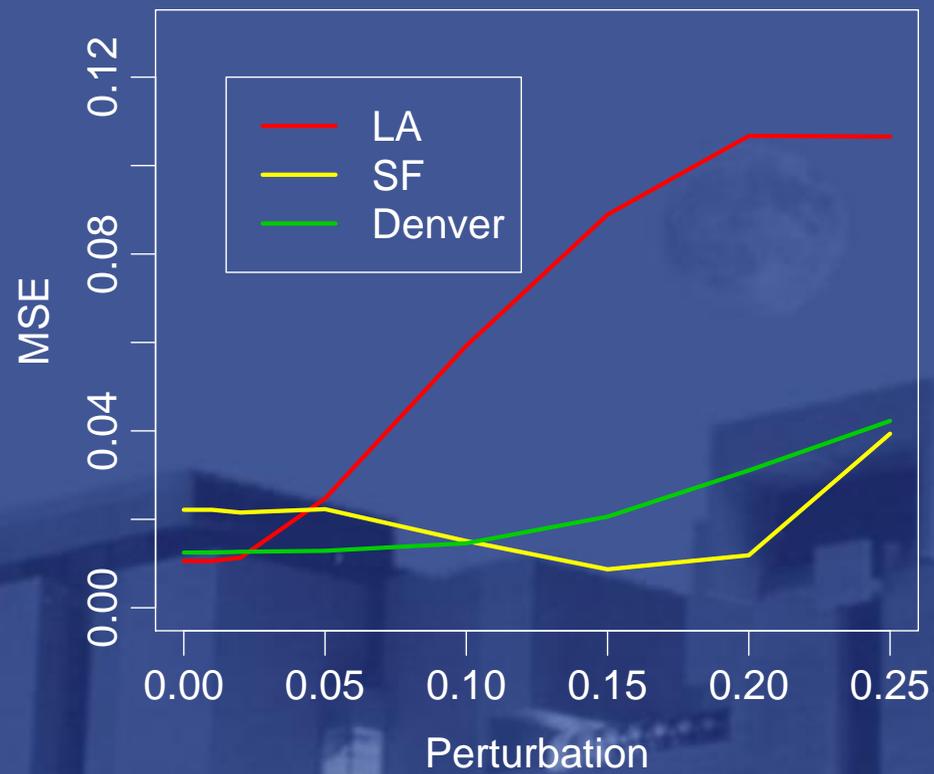
Measurement error



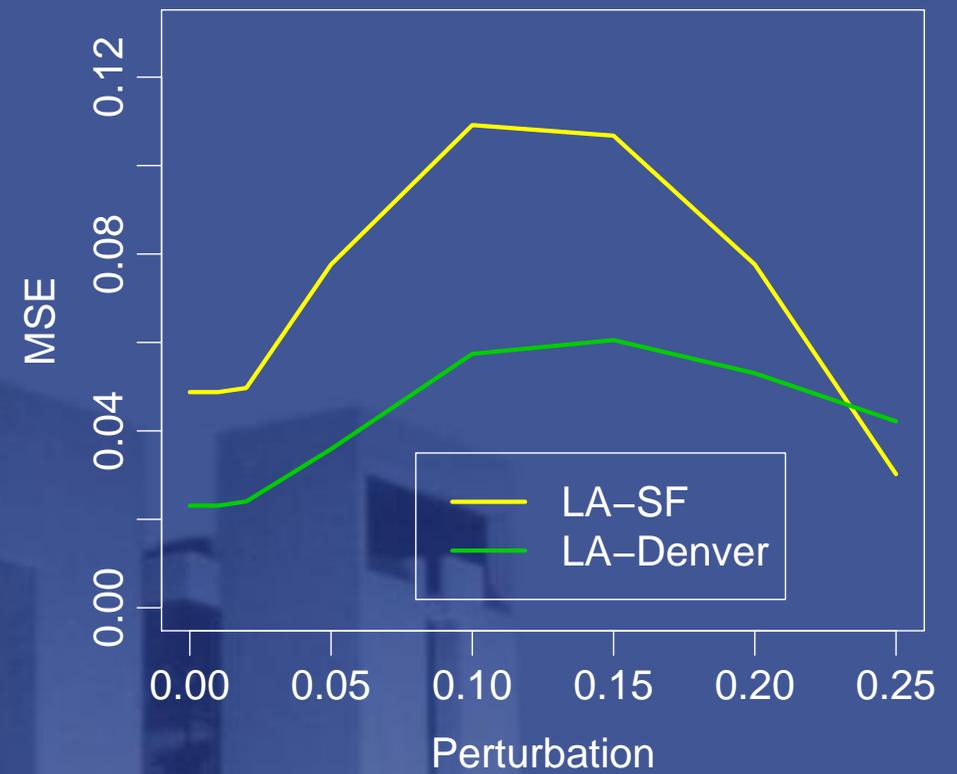
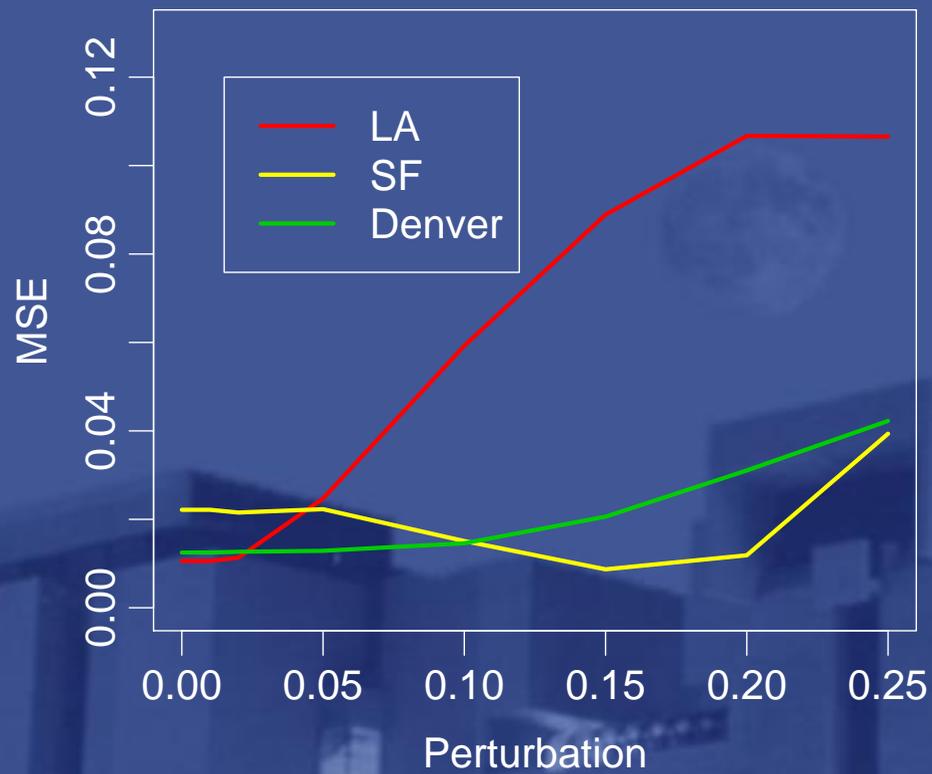
Measurement error and wind perturbation



Sensitivity Analysis: Results



Sensitivity Analysis: Results



Sensitivity Analysis

Regression model

$$\mathbf{y} = \mathbf{K}\boldsymbol{\beta} + \boldsymbol{\varepsilon}$$

$$\mathbf{y} = \mathbf{K}_{\text{pert}}\boldsymbol{\beta}_{\text{pert}} + \boldsymbol{\varepsilon}$$

where

\mathbf{y} concentration anomaly after 4 days

\mathbf{K} , \mathbf{K}_{pert} true and perturbed transport operator

$\boldsymbol{\varepsilon}$ measurement error

The mean squared error is

$$\text{MSE} = \sigma^2(\mathbf{K}^T\mathbf{K})^{-1}$$

$$\text{MSE} = \sigma^2(\mathbf{K}_{\text{pert}}^T\mathbf{K}_{\text{pert}})^{-1} + (\mathbf{M} - \mathbf{I})\boldsymbol{\beta}\boldsymbol{\beta}^T(\mathbf{I} - \mathbf{M}^T)$$

with $\mathbf{M} = (\mathbf{K}_{\text{pert}}^T\mathbf{K}_{\text{pert}})^{-1}\mathbf{K}_{\text{pert}}^T\mathbf{K}$.